NOW?

Thermoluminescence characteristics of USGS standard basaltic rock BCR-1

K. S. V. NAMBI, B. D. BHASIN AND V. N. BAPAT Heald: Physics Division, Bhabha Atomic Research Centre, Bombay-400 085 (India) (Received 15 August 1977)

Thermoluminance (TL) investigations on lunar and meteorite sampks are on the increase and it is necessary for the laboratories to work with some reference standard sample of similar behabiour before taking up actual measurements on precious samples of limited availability. The **different rock** sampIes of the United States geological survey standards can fill this need and measurements made on BCR-I, a standard basalt of the Columbia river bed are presented here.

EXPERIMENTAL

The TL instrument ussd is described elsewhere'; however, instead of the motor-driven autovariac for the heater voltage supply, an indigenously made temperature proqalnm **er was used in the present investigations. Ail TL measurements were normal&d to 5 rnzm wei@ of the sample on the TL heater pan. The TL emission** spectra were determined by recording monochromatic TL glow curves through band pass filters[?].

RESULTS AND DISCUSSION

The natural TL (NTL) glow curve of the virgin sample and that obtained after artificial gamma irradiation in the laboratory (ATL) are shown in Figs. 1 and 2. It is clearly seen that the TL glow curve is characterised by three peaks at $170^{\circ}C$ (I), 283 $^{\circ}$ C (II) and 377 $^{\circ}$ C (III) with the first peak having completely decayed in the NTL.

From the **known U, Th** and K concentrations in BCR-I, it can be computed that the **total beta, gamma annual self-irradiation** rate is of the order of 300 mR yr^{-1} . (The alpha irradiation can be *negid as* its **TL effeaiveness is very low_) The TL cAbration obtained by incremental artificial gamma irradiations in the laboratory** reveals that the NTL exhibited by BCR-I is equivalent to about 4×10^4 R of beta, gamma irradiation. This leads to a TL age estimate (upper limit, as the cosmic **component of the irradiation is neglected) of about** 1.4×10^5 **years which is quite low**

Fig. 1. NTL glowcurve of USGS standard basalt BCR-I: A, NTL curve; B, thermal background. $A - B = NTL$ signal. Heating rate 300°C min⁻¹.

Fig. 2. ATL glowcurve of USGS standard basalt BCR-L Gamma exposure given = 8.6×10^5 R. Heating rate 300°C min⁻¹.

compared to 20-30 million years, known² to be the age of BCR-I. It is also known³ that, in dating up to about 50 million years, an NTL peak at more than 300°C coupled with an activation energy nearer to about 2 eV is necessary; otherwise the peak is drained considerably even at 20°C. Although the NTL peak of BCR-I is at a temperature well above 300°C, the low TL age obtained may be indicative of a rather shallow trap $(< 2$ eV) with a mean life much shorter than the actual geological age involved. Experimental determination of the activation energy by the "initial rise method" has indeed yielded values of 0.73 eV and 1.42 eV respectively for the peaks II and III which constitute the NTL.

The TL emission spectra are presented in Fig. 3 and it is seen that peak I is emitted predominantly in the 500-560 nm region while peaks II and III (corresponding

Fig. 3. TL emission spectra of USGS standard basalt BCR-L

to NTL) are emitted predominantly in the 350–400 nm region. The bands could not **ix ~11 rc;okd becase of low levels of emission and the use of band pass filters to register the spectrum. However, these bands coincide remarkabIy with the emissions** of Mn^{2+} and Ce^{3+} ions which are well known TL emitters (at $\sim 1\%$ by weight and trace quantities respectively) in a variety of host lattices such as CaF₂, CaSO₄, CaCO₃, CaO, CaS etc. Available analytical reports⁴ reveal-that BCR-I has about **0.2% by weight of Mn and among the rare earths Ce is present in the highest concentrafion range of around 50 ppm_ The predominance of the manganese emission at** lower temperatures and cerium emission at higher temperatures lends further support **to the conclusion that Cc and Mn impurities cause the TL emission in BCR-I-**

ACKNOWLEDGEMENT

Thanks are due to Dr. M. Sankardas, Head, Analytical Division for making available the USGS standard,

REFERENCES

1 K. S. V. Nambi, V. N. Bapat and B. D. Bhasin, in N. Bhandari and M. N. Rao (Eds.), Farther Advances in Lunar Research: Luna 16 and 20 Samples, Indian National Science Academy, 1974, p. 111. $\sim 10^{-11}$

- 2 R. C. Newcombe, in U.S. Geological Survey Research Professional Paper, Series 424 B, 1961, p. 213-215.
- 3 G. Bonfiglioli, in D. J. McDougall (Ed.), Thermoluminescence of Geological Materials, Academic Press, 1968, p. 15-24.
- 4 M. Fleischer and F. J. Flanagen, Geochim. Cosmochim. Acta, 33 (1969) 55, 81.

÷.

 $\mathcal{L}^{\mathcal{L}}$, where $\mathcal{L}^{\mathcal{L}}$ and $\mathcal{L}^{\mathcal{L}}$

 ~ 10

i.

 $\mathcal{A}=\mathcal{A}$

 ~ 10

 $\label{eq:2.1} \mathcal{L}(\mathcal{L}^{\mathcal{L}}) = \mathcal{L}(\mathcal{L}^{\mathcal{L}}) = \mathcal{L}(\mathcal{L}^{\mathcal{L}}) = \mathcal{L}(\mathcal{L}^{\mathcal{L}})$

 ~ 1000 km s $^{-1}$

 \mathbb{R}^n

 \mathcal{L}